



I. Introduction

Campobello 7.5-minute quadrangle (CQ) is located on the Six Mile thrust sheet within the Tugaloo terrane of the Inner Piedmont. Geologic mapping of fifteen 7.5-minute quadrangles in the Western Inner Piedmont of South Carolina by Furman University and the South Carolina Geological Survey personnel (1995-2010) has investigated the polyphase fold deformation, faulting, and metamorphic history of the region. Our mapping traces these structures into CQ.



West side of Little Mountain (Collins Mountain). This prominent ridge in southeast Campobello quadrangle has 60 m relief and is the highest peak in the area. View to the east.



View to the northeast across peach orchards and rolling pasture land typical of northern Campobello quadrangle. Blue Ridge Front visible in the distance. View to northeast. Each year, South Carolina grows more than 200 million pounds of peaches. The state is the No. 2 peach producer in the nation, behind California. Georgia is No. 3.

II. Rock Type

Tallulah Falls Formation gneiss (Neoproterozoic) Interlayered biotite quartz feldspar gneiss, hornblende quartz feldspar gneiss, amphibolite, amphibole quartz feldspar gneiss, calc-silicate gneiss, and garnet biotite quartz feldspar gneiss. The different rock types are interlayered within CQ. The gneiss unit underlies ca. 95% of the quadrangle area.



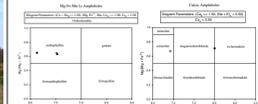
Schist with tight chevron folds (F2) that deform cm scale isoclinal folds (F3). Exposure lies north of the Beaverdam fault and is structurally on the north limb of an F2 synform. View to south.



Photomicrograph (FOV 2mm) of hornblende-biotite-quartz-feldspar gneiss displaying hornblende (H), albite-twinned plagioclase (P), and titanite (T) (high relief). Crossed polars.



The hornblende metagabbro body is surrounded by Tallulah Falls gneiss, but the contacts of the body are not exposed. Based on limited field data we tentatively suggest the hornblende metagabbro lies in the core of an F2 synform (see cross section D-D'). If true, it is probably a fairly thin intrusive body. In addition, the hornblende metagabbro outcrop area between the North and Holly Springs faults may be a horst.



Slabbed coarse-grained hornblende metagabbro specimen with relict igneous texture (?) cut by thin amphibolite dikes (?) (left and right margins). Photomicrograph (FOV 2mm) of recrystallized hornblende metagabbro. Hornblende (H) surrounds epidote (E) grains in the center of the image.

Tallulah Falls Formation schist (Neoproterozoic) Interlayered schist and schistose muscovite-biotite gneiss.

The original stratigraphic order of metamorphic units in CQ is obscured by polyphase folding and faulting. Schists now lie structurally above the Tallulah Falls Formation gneiss. Schist and gneiss in the map unit have either biotite or muscovite or both present. Common lithologies include garnet (up to 5 mm)-mica schist; mica schist interlayered with pegmatite; fine-crystalline, limonitic-stained and beige-weathered, leucocratic, muscovite-sillimanite schist with sheared quartz lenses, 1-3mm thick and 0.5-2cm long; and a fine-crystalline, darker limonitic stained, muscovite schist.

With a decrease in modal mica and an increase in quartz and feldspar, schist grades into gneiss. The gneiss varieties include muscovite-biotite-garnet-quartz gneiss, with garnet comprising up to 40% of the rock composition and muscovite-biotite-quartz-plagioclase (oligoclase) gneiss.



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Holly Springs Hornblende Metagabbro A hornblende metagabbro body (3 x 0.5 km) is located between Lyman Lake and Holly Springs, SC. It is partly truncated and offset by east-northeast faults. Metagabbro is largely hornblende and plagioclase, the latter altering to epidote and scapolite in thin section. Some textures suggest the rock retains an original igneous texture. In float samples, zones of foliated amphibolite several cm wide with sharp contacts may represent original thin basaltic dikes in the gabbro.

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Silicified cataclastic rocks Microbreccia, cataclastite, and syntaxial quartz veins ("comb quartz") form narrow zones of outcrop, boulders, and resistant float. Discontinuous cataclastic rock bodies up to 1.3 km in length and trending N 60°-70°E adjacent to the Pax Mountain Fault have been traced across the entire Campobello quadrangle. On a ridge 1.6 km northeast of Liberty Church, two main zones of outcrops and boulders of microbreccia and comb quartz are 10m and 25m long. In general, linear cataclastic rock bodies are a few meters long and up to 3m high.



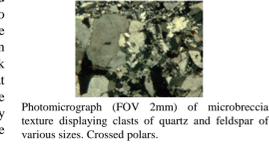
Microbreccia body 12 m wide of Pax Mountain fault zone and enclosing leucocratic biotite quartz-feldspar gneiss is exposed along Rt. 176 in Campobello, SC. View to northeast.



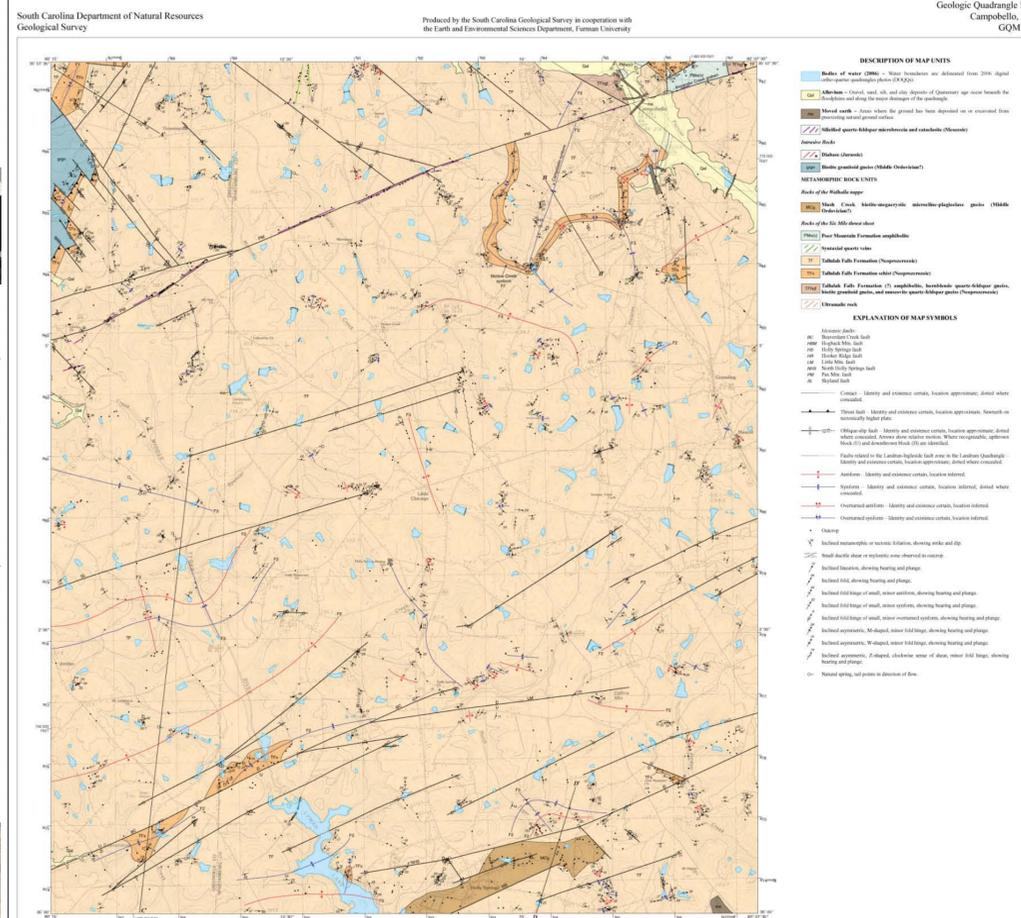
Thin, criss-crossing microbreccia and "comb quartz" veins (striking N40-80 E and N 18 W) are exposed over a 17m zone south of the 12m-wide Campobello microbreccia body. Pinkish enclosing rock is leucocratic biotite quartz-feldspar gneiss.



Continuous zone of microbreccia exposures and float (32 ft long) trends N45°E along Pax Mountain fault. Ridge 1.6 km northeast of Liberty Church.

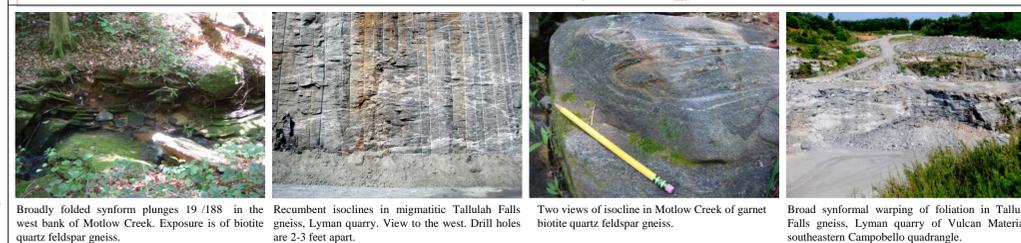
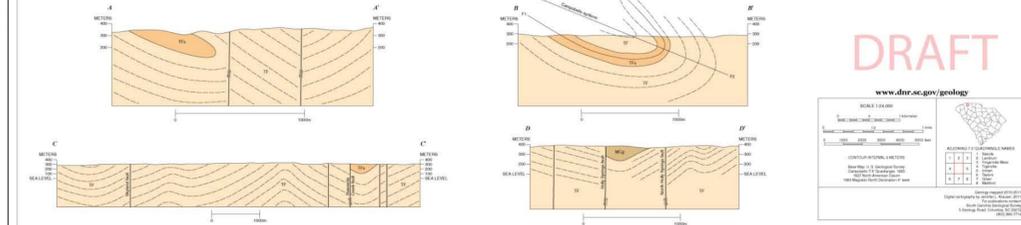


Photomicrograph (FOV 2mm) of microbreccia texture displaying clasts of quartz and feldspar of various sizes. Crossed polars.

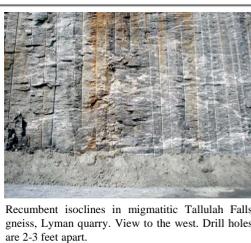


GEOLOGIC MAP OF THE CAMPOBELLO QUADRANGLE, GREENVILLE AND SPARTANBURG COUNTIES, SOUTH CAROLINA

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Broadly folded synform plunges 19/188 in the west bank of Motlow Creek. Exposure is of biotite quartz feldspar gneiss.



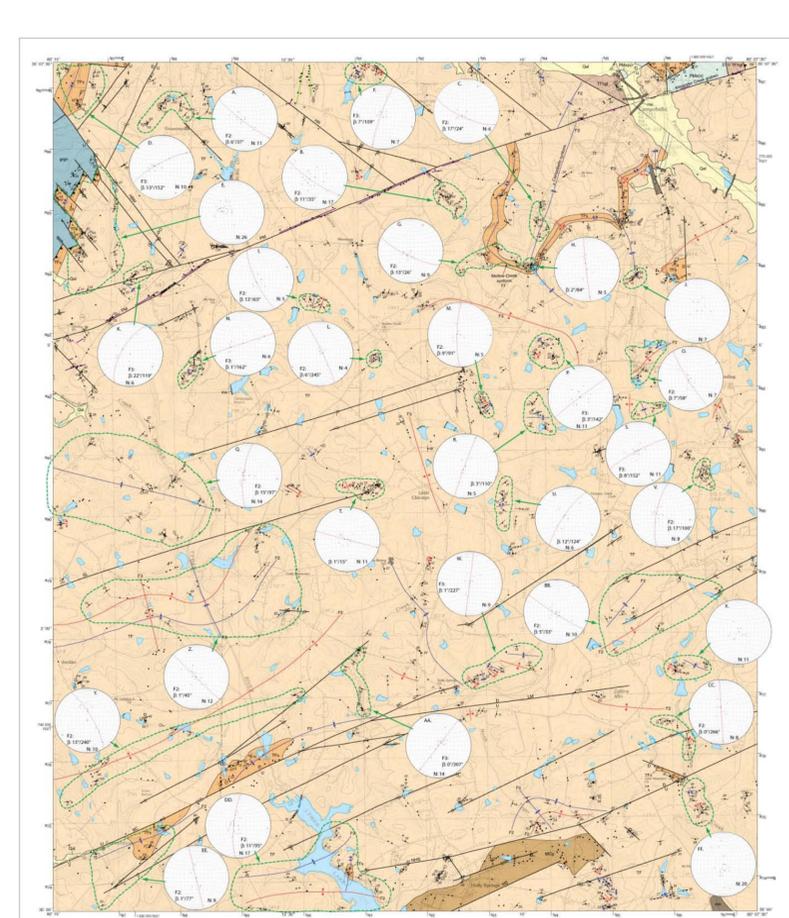
Recumbent isoclinal folds in migmatic Tallulah Falls gneiss, Lyman quarry. View to the west. Drill holes are 2-3 feet apart.



Two views of isoclinal in Motlow Creek of garnet biotite quartz feldspar gneiss.



Broad synformal warping of foliation in Tallulah Falls gneiss, Lyman quarry of Vulcan Materials, southeastern Campobello quadrangle.



SOUTHERN HEMISPHERE, EQUAL AREA STEREOPLOTS OF POLES TO FOLIATION SURFACES ARE USED TO DETERMINE STATISTICAL FOLD HINGE ORIENTATIONS (BETA) OF F2 AND F3 FOLDS

III. Cross Sections

A-A': In the northwest of CQ, Tallulah Falls schist lies structurally above Tallulah Falls gneiss in the nose of an overturned synform. The synform is truncated by a vertical fault. B-B': In the northeast of CQ and to the southeast of the Pax Mountain fault, schist forms a long, sinuous belt which has been deformed by F1, F2, and F3 generation folds. The cross section shows part of this schist belt, which is structurally an overturned synform and consists of interlayered Tallulah Falls gneiss and schist. The earlier F1 isoclinal fold was later refolded by F2 folding. C-C': The Beaverdam Creek Fault, trending N55 E, in southeastern CQ truncates the schist unit, which occurs in the nose of a synform. The gneiss unit displays broad synforms and antiforms and is faulted by vertical faults. D-D': Hornblende Metagabbro body which is located in the southeast of CQ between Lyman Lake and Holly Springs, SC.

IV. Metamorphic History

Although metamorphism peaked in the sillimanite zone of the upper amphibolite facies, we observe a range of metamorphic conditions in the rocks of the CQ. Most notably, chlorite is developed in most of the amphibole quartz feldspar gneiss lithologies in which the amphibole is altered to chlorite. Microprobe data indicate that several amphiboles exist, which indicates a range of temperatures and pressures from the greenschist facies to the upper amphibolite facies. On the geologic map of Campobello quadrangle a trend appears for four of the locations of retrograded amphibole gneiss at N50 E and is a similar strike with nearby faults mapped in Campobello quadrangle. This similarity in trends suggests that the faults might have served as avenues for fluids, which then produced the retrograded mineralogy in localized regions.

V. Deformation: Faults and Folds

Folds in Campobello quadrangle F1 - Isoclinal folds are found at all scales. The F1 designation is used for the oldest folds we observed in CQ. Foliation and compositional layering are affected by F1 folding. Hence it is unknown whether the compositional layering in gneiss is in fact axial planar to a fold set older than our F1 set. Macroscopic F1 folds were probably overturned (recumbent?) originally, with a northwest-vergent consistent with Griffin's nappe emplacement concepts for the Inner Piedmont (Wetmore, C.C., and Griffin, V.S., 1977; Griffin, V.S., 1971; Griffin, V.S., 1967). Aligned sillimanite needles may have formed during F1 deformation. In rock exposures a transposition foliation has been observed.



Small isoclinal fold in biotite gneiss, Lyman quarry.

F2 - Inclined to overturned macroscopic and mesoscopic F2 folds deform earlier folds.



Overturned F2 synform in interlayered biotite quartz-feldspar gneiss and amphibolite, Motlow Creek. Hammer lies on the gently dipping upright fold limb; to the right, layers abruptly bend and are overturned in the background at the narrow ledge crossing the creek. F2 fold hinge plunges to the upper left at 30°/N59 E.

F3 - Inclined to overturned macroscopic and mesoscopic F3 folds deform earlier and F2 folds.

F4 and F5 - Upright gentle folds. Not known from macroscopic map patterns



Pavement of biotite quartz feldspar gneiss along Meadow Creek, displaying a Type 1 interference fold pattern due to polyphase deformation. The white tape measures mark the axial traces of two gentle synforms which trend northwest and northeast.

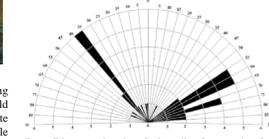
Polyphase folding in Campobello quadrangle

Campobello quadrangle has experienced five generations of folding resulting in polyphase deformation. The Motlow Creek F1 isoclinal folds in the schist unit located in the northeast of CQ has been deformed by the macroscopic Campobello F2 synform. Other F2 folds are seen as tight chevron folds found in schist and have deformed F1 aligned sillimanite needles, as evident from hooks that are visible in chevron folded schist. F3 macroscopic inclined to overturned folds trend northwest and warp F1 and F2 folds. This can be seen clearly in the Campobello synform (F2), which is cross folded by F2 folds, F2 and F3 together create Type-2 interference folding. The final two generations of folding are mesoscopic F4 and F5, two generations of gentle folding that appear 90° to each other in stream pavements creating Type-1 interference folding.

Faults in Campobello quadrangle

The chronology of faulting in CQ is northeast (oldest), northerly, northwest, and east-northeast to easterly faults. Dominant fault sets consist of older and shorter N35 - 40 W striking faults and younger but longer N55 - 75 E striking faults.

Zones of microbreccia boulders lie along the Pax Mountain fault and trend N30 - 60 E, with individual sets trending N30 E, N40 - 50 E, N55 E and N60 E. In an exposure of microbreccia in a roadside ditch at GPS location 387668/3883602, near Morrow Road and less than 10m north of the trace of the Pax Mountain fault, extensional veins up to 1ft long were oriented N25 E 68 NW, N50 E 60 NW, and N55 E 82 NW. Northeast of this microbreccia exposure and still less than 10m from the Pax Mountain fault, the orientation of syntaxial veins was measured as N50 E, E-W, and N40 - 70 E. Two separate, north-trending syntaxial quartz veins (0.5-1 km long) lie 2 km northeast of Gowensville.



Rose Diagram showing fault strike frequencies for CQ



Tape and brunton map of pavement. Foliation form lines are shown. A structural basin at X (GPS location 390978/3876805) is formed by two cross-cutting, gentle synforms: a N53 W fold axial trace labeled S1-S1' and a N37 E axial trace labeled S2-S2'.



Synoptic stereonet plot analysis of 24 hinge orientations of F2 and F3 hinge data display two hinge orientations. The hinges for F2, in black range in azimuths of 025° to 100° and 240° to 270°. The hinges for F3, in purple range in azimuths of 100° to 170° with just one point plunging northwest. Our interpretation of this distribution is F2 folds being folded about F3 with a gentle plunge southeast.

VI. Summary

- The metamorphic peak is in the sillimanite zone of the upper amphibolite facies for the rocks in Campobello quadrangle. Localized retrograde fluids along faults trending N50 E metamorphosed rocks to greenschist facies, leading to chlorite replacement of amphiboles. Campobello quadrangle has experienced multiple deformational events resulting highly faulted and polyphase deformed geologic units. The fold chronology consists of five fold generations. F1 mesoscopic and macroscopic isoclinal folds, F2 inclined to overturned mesoscopic chevrons and macroscopic folds, F3 gently inclined to overturned mesoscopic and macroscopic folds, and F4 and F5 gentle folds. The faulting chronology is (oldest-youngest): northeast, northerly, northwest, and east-northeast to northerly faults. Dominant fault trends are N35°-40°W and N55°-75°E and faults are assumed to be vertical.

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